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Rock mechanics modelling of rock mass properties – summary of primary data

Preliminary site description Laxemar subarea – version 1.2

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May 2006

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This report concerns a study which was conducted for SKB. The conclusions and viewpoints presented in the report are those of the authors and do not necessarily coincide with those of the client.

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Abstract

The results presented in this report are the summary of the primary data for the Laxemar Site Descriptive Modelling version 1.2. At this stage, laboratory tests on intact rock and fracture samples from borehole KSH01A, KSH02A, KAV01 (already considered in Simpevarp SDM version 1.2) and borehole KLX02 and KLX04 were available.

Concerning the mechanical properties of the intact rock, the rock type "granite to quartz monzodiorite" or "Ävrö granite" (code 501044) was tested for the first time within the frame of the site descriptive modelling. The average uniaxial compressive strength and Young's modulus of the granite to quartz to monzodiorite are 192 MPa and 72 GPa, respectively. The crack initiation stress is observed to be 0.5 times the uniaxial compressive strength for the same rock type. Non negligible differences are observed between the statistics of the mechanical properties of the granite to quartz monzodiorite in borehole KLX02 and KLX04.

The available data on rock fractures were analysed to determine the mechanical properties of the different fracture sets at the site (based on tilt test results) and to determine systematic differences between the results obtained with different sample preparation techniques (based on direct shear tests).

The tilt tests show that there are not significant differences of the mechanical properties due to the fracture orientation. Thus, all fracture sets seem to have the same strength and deformability. The average peak friction angle for the Coulomb's Criterion of the fracture sets varies between 33.6° and 34.1°, while the average cohesion ranges between 0.46 and 0.52 MPa, respectively. The average of the Coulomb's residual cohesion and friction angle vary in the ranges 28.0°–29.2° and 0.40–0.45 MPa, respectively. The only significant difference could be observed on the average cohesion between fracture set S A and S d.

The direct shear tests show that the mechanical properties obtained from the laboratory tests very much depend on the sample preparation technique and size of the steel ring holders. The tests performed with concrete mould or with large steel ring holders (SP results) present larger deformability compared to the tests performed with epoxy resin mould and small steel ring holders (NGI results). The normal and shear stiffness of the fractures obtained by NGI are on average 608 and 21 MPa/mm, while their standard deviation is 394 and 9 MPa/mm, respectively. For comparison, the normal and shear stiffness of the fractures obtained by SP are on average 135–237 and 29–41 MPa/mm, respectively. Due to the limitation in amount of data from direct shear tests, a determination of the properties for each fracture set was not possible.

Sammanfattning

De resultat som presenteras i denna rapport är en sammanfattning av primärdata för Laxemar platsbeskrivande modell version 1.2. Laboratorietester av intakt berg och sprickor innefattar prover från borrhål KSH01A, KSH02A, KAV01 (tidigare redovisade i Simpevarp modell version 1.2) samt borrhål KLX02 och KLX04.

Beträffande de mekaniska egenskaperna hos intakt berg testades bergarten "granit till kvarts-monzodiorit" eller "Ävrö granit" (kod 501044) för första gången inom ramen för Laxemars platsbeskrivande modellering. Medelvärden för den enaxiella tryckhållfastheten och elasticitets modulen hos granit till kvartsmonzodiorit är 192 MPa respektive 72 GPa. Den s k "crack initiation stress" har observerats vara 0,5 gånger den enaxiella tryckhållfastheten för samma bergart. Man kan se icke-negligerbara skillnader i de bergmekaniska testresultaten för granit till kvartsmonzodiorit mellan borrhål KLX02 och borrhål KLX04.

Tillgänglig data från tester på bergsprickor har analyserats för att undersöka de mekaniska egenskaperna hos olika sprickgrupper i Laxemar/Simpevarp (baserat på tilttestresultat) och för att undersöka systematiska skillnader mellan uppnådda resultat vid olika typer av provpreparering (baserat på direkta skjuvtester).

Tilttesterna visar att det inte finns någon signifikant skillnad hos de mekaniska egenskaperna med avseende på sprickorientering. Sålunda verkar alla sprickgrupper oavsett orientering ha samma hållfasthet- och deformationsegenskaper. För Coulombs brottskriterium har sprickgruppernas maximala pikfriktionsvinkel medelvärden mellan 33,6° och 34,1° medan maximala pik-kohesionen har medelvärden mellan 0,46 och 0,52 MPa. Medelvärdet för de residuala friktionsvinkeln och kohesionen varierar mellan 28,0° och 29,2° respektive 0,40 och 0,45 MPa. Den enda signifikanta skillnaden mellan sprickgrupper observerades för kohesionen mellan sprickgrupp S_A och S_d.

Direkta skjuvtester visar att de mekaniska egenskaperna till stor del beror på provprepareringen och storleken på stålringshållarna. Tester gjorda med gjuten betong eller epoxy och stora stålringshållare (SP-resultat) ger lägre hållfasthet och större deformerbarhet, jämfört med tester gjorda med gjuten epoxy och små stålringshållare (NGI-resultat). Medelvärdet för sprickornas normal- och skjuvstyvhet är 608 respektive 21 MPa/mm och standardavvikelsen är 394 respektive 9 MPa/mm, för NGI-resultaten. Som jämförelse, normal- och skjuv-styvheten beräknat från SP-resultat har medelvärde varierande mellan 135 och 237, och respektive 29 och 41 MPa/mm. Den tillgängliga datamängden från direkta skjuvtester tillät inte en bestämning av egenskaperna för varje sprickgrupp.

Contents

1 Introduction

Table 1-1 lists the number of intact rock samples tested in laboratory for the determination of the strength and deformability. Uniaxial, triaxial and indirect tensile tests were performed on these samples. The results of the tests are summarised in Chapter 2. Table 1-2, on the other hand, lists the number of natural fracture samples tested in laboratory. The tests performed were direct shear and tilt tests. The results of the tests on fractures are summarised in Chapter 3. A map of the site with the location of the five boreholes where the samples were collected is shown in Figure 1-1.

Boreholes KSH01A, KSH02A and KAV01 are located in the Simpevarp site investigation area, and boreholes KLX02 and KLX04 are located at the Laxemar site.

Borehole	Rock type	Indirect tensile tests	Uniaxial tests	Triaxial tests
KSH01A	Quartz monzonite to monzodiorite ¹	20 [2 ²]	10	8 [3 ²⁾ , 1 ³]
	Fine-grained dioritoid ¹	$20 [6^2]$	10 $[4^{2}]$	4 $[1^2]$
KSH02A	Fine-grained dioritoid ¹	12 [2 ²]	$5 [1^2]$	$5[2^2]$
KLX02	Granite to quartz monzodiorite	30	15	
KLX04	Granite to quartz monzodiorite	30	15	14

Table 1‑1. Summary of tests performed on intact rock samples from borehole KSH01A, KSH02A, KLX02 and KLX04.

1) Data already reported in Simpevarp SDM version 1.2 /SKB 2005/.

2) Samples with sealed fractures.

³⁾ Samples with intrusion of fine to medium grained granite.

Table 1‑2. Summary of the results of tests performed on fracture samples from borehole KSH01A, KSH02A, KAV01, KLX02, and KLX04.

¹⁾ 3 levels of normal stress: 0.5, 5 and 20 MPa.

 $2)$ 5 samples from depth $> 1,000$ m.

Figure 1‑1. Overview map of core-drilled and percussion-drilled boreholes in the Laxemar and Simpevarp subareas.

2 Intact rock

In this Chapter, the results of the uniaxial compressive strength tests on intact rock samples are summarised independently and together with the results of the triaxial compressive strength tests. The rock types represented are: fine-grained dioritoid, quartz monzonite to monzodiorite and granite to quartz monzodiorite. The SICADA codes for these rock types are listed in Table 2-1.

The laboratory results are reported in:

- /Jacobsson 2004abcd/ for the uniaxial compression tests.
- /Jacobsson 2004fgh/ for the triaxial compression tests.
- /Jacobsson 2004ijkl/ for the indirect tensile tests.

In the following sections, only the updated set of mechanical properties for Laxemar Site Descriptive Model 1.2 compared to Simpevarp SDM 1.2 are provided. For all the properties that are unchanged, please refer to /SKB 2005/.

Table 2‑1. List of the rock types and their SICADA code for the samples tested in laboratory for Laxemar SDM 1.2.

Rock type	SICADA code
Fine-grained dioritoid (metavolcanite, volcanite)	501030
Quartz monzonite to monzodiorite (equigranular to weakly porphyritic)	501036
Granite to quartz monzodiorite (generally porphyritic) - Ävrö granite	501044

2.1 Uniaxial compressive strength

Laboratory tests of uniaxial compressive strength UCSi were carried out at the SP Laboratory (Swedish National Testing and Research Institute) on samples from borehole KSH01A, KSH02A, KLX02 and KLX04 /Jacobsson 2004abcd/. The results are given per rock type by means of the mean value and the standard deviation in Table 2-2. The statistical description is completed with the minimum, maximum and most frequently occurring values.

In Figure 2-1, earlier results reported in Simpevarp SDM 1.1 /SKB 2004/ are also shown for comparison with the new frequency distributions of the uniaxial compressive strength in Figure 2-2.

It is worth it to notice that, compared to Simpevarp SDM 1.2, only samples of granite to monzodiorite were added to the sample set. For comparison, some statistics of the uniaxial compressive strength for borehole KLX02 and KLX04 are listed in Table 2-3. Considering the fact that many samples are available for each borehole (15), the difference between the

calculated statistics are significant (about 11% and 15% for the mean value and the standard deviation, respectively). In this table, the statistics for the samples containing sealed fractures are also reported.

Results were reported by the HUT Laboratory (Helsinki University of Technology) on 5 samples of quartz monzonite to monzodiorite taken from borehole KSH01A /Eloranta 2004a/. These results were not included in Table 2-2.

1) Data already reported in Simpevarp SDM version 1.2 /SKB 2005/.

²⁾ Results are not reported in SICADA for 1 test on a sample from KLX02.

¹⁾ Results are not reported in SICADA for 1 test on a sample from KLX02.

Figure 2‑1. Frequency distributions of the uniaxial compressive strength of all samples available from Äspö and CLAB. The rock types were not specified in the sources.

2.1.1 Crack initiation stress

The crack initiation stress σ_{ci} was calculated according to /Martin and Chandler 1994/. For samples of granite to monzodiorite, the values in Table 2-4 and the frequency distribution in Figure 2-3 apply. As for the uniaxial compressive strength, differences in the statistics are observed between samples from KLX02 and KLX04 (Table 2-5).

Table 2‑4. The crack initiation stress σ**ci from uniaxial compressive tests performed on intact rock samples from KLX02 and KLX04.**

Rock type	Number of	Minimum $\sigma_{\scriptscriptstyle \rm ci}$	Mean $\sigma_{\scriptscriptstyle{\text{ci}}}$	Frequent σ_{ci}	Maximum σ_{ci}	$\sigma_{\rm el}$'s Standard
	samples	[MPa]	IMPal	IMPal	[MPa]	deviation [MPa]
Granite to monzodiorite	29^{1}	32.0	95.8	100.0	126.5	20.7

¹⁾ Results are not reported in SICADA for 1 test on a sample from KLX02.

Table 2‑5. Comparison of the crack initiation stress from uniaxial compressive tests obtained for samples from borehole KLX02 and KLX04.

¹⁾ Results are not reported in SICADA for 1 test on a sample from KLX02.

Uniaxial compressive strength [MPa]

Figure 2‑2. Frequency distributions of the uniaxial compressive strength of the samples of finegrained dioritoid, quartz monzonite to monzodiorite and granite to quartz monzodiorite from borehole KSH01A, KSH02A, KLX02 and KLX04.

Figure 2‑3. Crack initiation stress for the granite to monzodiorite from the uniaxial compression testing of samples from borehole KLX02 and KLX04.

2.1.2 Correlation between uniaxial compressive strength and crack initiation stress

The values of the crack initiation stress in Section 2.1.1 can be plotted against the uniaxial compressive strength in Section 2.1 so that Figure 2-4 can be obtained. This figure shows that there is a clear correlation between the two values that can be approximated with a linear trend. The slope of such trend is 0.5.

Figure 2‑4. Uniaxial compressive strength versus crack initiation stress for the granite to monzodiorite of the samples from borehole KLX02 and KLX04.

2.2 Triaxial compressive strength

Laboratory tests of triaxial compressive strength were carried out at the SP Laboratory (Swedish National Testing and Research Institute) on samples from borehole KSH01A, KSH02A and KLX04 /Jacobsson 2004fgh/. In the following analyses, the results of triaxial testing are considered together with the results of uniaxial testing.

For each main rock type (fine-grained dioritoid, quartz monzonite to monzodiorite, granite to quartz monzodiorite), the triaxial results were analysed together with the correspondent results of the uniaxial compressive tests. The laboratory results on intact rock samples were interpolated with the Hoek and Brown's Failure Criterion /Hoek et al. 2002/.

$$
\sigma_1 = \sigma_3 + UCS_T \left(m_i \frac{\sigma_3}{UCS_T} + 1 \right)^{0.5}
$$
 (1)

where σ' ₁ and σ' ₃ are the maximum and minimum principal stress and m_i is a strength parameter typical for each rock type. UCSi_{T} is obtained by matching the uniaxial and triaxial test results and thus slightly differs from UCSi in Section 2.1.

When analysing the laboratory results, the intact rock parameters in Table 2-6 are obtained. Although obtained in a slightly different way, the results of the UCSi are in rather good agreement with the values in obtained on uniaxial tests only (Table 2-2).

The Coulomb's linear approximations of the Hoek and Brown's Criterion were also calculated for a certain stress interval (0 to 15 MPa, Table 2-7). These linear approximations are shown in Figure 2-5, Figure 2-6 and Figure 2-7 for the fine-grained dioritoid, quartz monzonite to monzodiorite, granite to quartz monzodiorite, respectively. The Hoek and Brown's Criterion also provides an estimation of the tensile strength of the intact rock that can be compared with the laboratory results in Section 2.3. The statistics for the samples containing sealed fractures are also reported.

Five samples of quartz monzonite to monzodiorite were also tested in triaxial compression conditions at the HUT Laboratory /Eloranta 2004b/ under confining pressures of 2, 7 and 10 MPa. These results are in good agreement with the SP Laboratory results but were not included in Table 2-6 and Table 2-7.

Table 2‑6. Parameters for the Hoek & Brown's Criterion based on the results of uniaxial and triaxial tests performed on intact rock samples from borehole KSH01A, KSH02A, KLX02 and KLX04.

1) Data reported in Simpevarp SDM version 1.2 /SKB 2005/.

Figure 2‑5. Hoek & Brown's and Coulomb's failure envelopes from uniaxial and triaxial tests for the fine-grained dioritoid. Data reported in Simpevarp SDM version 1.2 /SKB 2005/.

Figure 2‑6. Hoek & Brown's and Coulomb's failure envelopes from uniaxial and triaxial tests for the samples of quartz monzonite to monzodiorite. Data reported in Simpevarp SDM version 1.2 /SKB 2005/.

Figure 2‑7. Hoek & Brown's and Coulomb's failure envelopes from uniaxial and triaxial tests for the samples of granite to quartz monzodiorite.

1) Data already reported in Simpevarp SDM version 1.2 /SKB 2005/.

The values of the cohesion and friction angle are obtained for a confinement stress between 0 and 15 MPa.

2.3 Indirect tensile strength

Indirect tensile tests were conducted on 143 core samples at the SP Laboratory (KSH01A, KSH02A, KLX02 and KLX04 /Jacobsson 2004ijkl/).

Table 2-8 contains the statistics of the test results for the fine-grained dioritoid, quartz monzonite to monzodiorite, granite to quartz monzodiorite, respectively. Figure 2-8 also shows the frequency distributions of the indirect tensile strength for each of the main rock types. Statistics for the samples containing sealed fractures are also reported.

Rock type	Number of samples	Minimum TS [MPa]	Mean TS [MPa]	Frequent TS MPal	Maximum TS [MPa]	TS Standard deviation [MPa]
Fine-grained dioritoid ¹	24	14	19	19	24	2
Quartz monzonite to monzodiorite ¹	18	12	18	17	24	4
Granite to quartz monzodiorite	60	9.3	13.0	13.1	16.4	1.5
Sealed fractures in intact rock ¹⁾	10	9	14	15	22	5

Table 2‑8. Summary of the results of indirect tensile tests performed on intact rock sampled from borehole KSH01A, KSH02A, KLX02 and KLX04.

1) Data already reported in Simpevarp SDM version 1.2 /SKB 2005/.

Five samples of quartz monzonite to monzodiorite were tested in indirect tensile conditions at the HUT Laboratory /Eloranta 2004c/. The results are in agreement with the SP Laboratory results but were not included in Table 2-8.

*Figure 2‑8. Frequency distribution of the indirect tensile strength of the samples of fine-grained dioritoid and quartz monzonite to monzodiorite. *) Data already reported in Simpevarp SDM version 1.2 /SKB 2005/.*

2.4 Young's modulus

Based on the stress-strain curves obtained for the tests reported in Section 2.1 and 2.2, the deformation modulus of the intact rock could be obtained.

2.4.1 Uniaxial loading

Table 2-9 presents a summary of the deformability results from the uniaxial compression tests. The same data are also shown as histogram for the granite to quartz monzodiorite in Figure 2-9. For the other rock types, no new laboratory results are available compared to Simpevarp SDM version 1.2 /SKB 2005/, therefore their statistics are unchanged in Table 2-9. The Young's modulus of the samples of intact fine-grained dioritoid is lower than for the samples containing sealed fractures.

Table 2‑9. Summary of the results of Young's modulus from uniaxial compressive tests performed on intact rock samples from borehole KSH01A, KSH02A, KLX02 and KLX04.

1) Data already reported in Simpevarp SDM version 1.2 /SKB 2005/.

²⁾ Results are not reported in SICADA for 1 test on a sample from KLX02.

Figure 2‑9. Frequency distributions of the Young's modulus of the granite to quartz monzodiorite from tests for Laxemar 1.2.

2.4.2 Triaxial loading

Even for the triaxial tests, all the new samples were taken from granite to quartz monzodiorite, therefore these are the new values in Table 2-10 compared to Simpevarp SDM version 1.2 /SKB 2005/. Also in this case, the Young's modulus of the samples of intact rock is lower than for the samples containing sealed fractures.

1) Data already reported in Simpevarp SDM version 1.2 /SKB 2005/.

2.5 Poisson's ratio

Also the Poisson's ratio can independently be obtained from the stress-strain curves of the tests reported in Section 2.1, uniaxial conditions, and Section 2.2, triaxial conditions, respectively.

2.5.1 Uniaxial loading

From the uniaxial compression tests, the statistics of the Poisson's ration in Table 2-11 can be obtained.

1) Data already reported in Simpevarp SDM version 1.2 /SKB 2005/.

²⁾ Results are not reported in SICADA for 1 test on a sample from KLX02.

2.5.2 Triaxial loading

From the triaxial compression tests, the statistics of the Poisson's ration in Table 2-11 can be obtained.

Table 2‑12. Summary of the results of Poisson's ratio from triaxial compressive tests performed on intact rock sampled from borehole KSH01A, KSH02A, KLX02 and KLX04.

1) Data already reported in Simpevarp SDM version 1.2 /SKB 2005/.

3 Natural fractures

The strength and deformability of the natural rock fractures was determined in two ways:

- 1) By means of tilt tests where shearing is induced by the self-weight of the upper block when the fracture is progressively tilted.
- 2) By means of direct shear tests where shearing is induced by actuators that apply a load perpendicular and parallel to the fracture plane. Three different types of shear test techniques were used; these are referred to as: Shear I, Shear II and Shear III. The main differences between the different shear test techniques are explained in Section 3.2.

The samples of natural fractures were taken from boreholes KAV01, KSH01A, KSH02A, KLX02, and KLX04.

Direct shear tests were performed on altogether 54 fracture samples. These were taken from boreholes KLX02 and KLX04 /Jacobsson 2004mn/, KAV01 /Jacobsson 2005o/, KSH01A /Jacobsson 2005p, Chryssanthakis 2004e/ and KSH02A /Jacobsson 2005q/. Seven fracture samples from borehole KSH01A were tested by NGI and 47 samples were tested by the Swedish National Testing and Research Institute (SP). To examine how the clamping of rock specimens in the laboratory test apparatus may influence test results, three different techniques were used. The main difference between the different methods is the casting material and size of the steel holders, which are used to hold the specimen in the shear test apparatus.

Tilt tests were performed on 157 fracture samples, which were taken from all boreholes (KAV01, KSH01A, KSH02A, KLX02, and KLX04) /Chryssanthakis 2003, Chryssanthakis 2004abcd/. All tilt tests were performed by the Norwegian Geological Institute Laboratory (NGI).

The laboratory results are evaluated in terms of several different rock mechanics parameters that describe fracture strength and deformability. These parameters are summarized and analysed in the following sections. It is of interest to analyze if the different fracture sets from the Simpevarp and Laxemar site have different fracture strength and deformability. Therefore the fractures were grouped into fracture sets at the site by matching the reported fracture depth with the BOREMAP records in SICADA according to the Descrete Fracture Network DFN model of the site /Hermansson et al. 2005/ (Table 3-1; see also Appendix 2). However, as explained above, the parameters have been derived by different laboratory test methods, which may entail systematic differences in results. The objective is therefore to distinguish whether significant differences can be found in laboratory data, in terms of: 1) fracture sets, and 2) laboratory test methods.

Table 3‑1. Orientation of the fracture sets at Simpevarp and Laxemar /Hermansson et al. 2005/.

Simpevarp			Laxemar		
Fracture set	Strike (right rule) [°]	Dip [°]	Fracture set	Strike (right rule) [°]	Dip [°]
S_A	158	86	S A	150	84
S B	280	90	S B	105	89
S_C	033	89	s c	022	86
S_d	183	28	S d	264	08
S f	063	66	S e	247	75

3.1 Tilt tests

Tilt tests were carried out on 157 samples from boreholes KAV01, KSH01A, KSH02, KLX02, and KLX04 /Chryssanthakis 2003, Chryssanthakis 2004abcd/. The tilt tests are designed to suit the fracture parameter determination according to /Barton and Bandis 1990/. The shear strength of the fracture is a function of the normal stress σ_n as:

$$
\tau = \sigma_n \tan \left[\Phi_b^{BB} + JRC \log \left(\frac{JCS}{\sigma_n} \right) \right]
$$
 (2)

JRC is Joint Roughness Coefficient that quantifies roughness, JCS is Joint Wall Compression Strength of the rock surfaces, and Φ_b^{BB} is basic friction angle on dry saw-cat surfaces, respectively. The residual friction angle Φ_r^{BB} is used instead of Φ_b^{BB} if the strength of wet surfaces is concerned. The index notation BB is used to emphasize that the parameters relate to the Barton-Bandis model, to differentiate them from parameters in the Mohr-Coulomb model, discussed later. /Barton and Bandis 1990/ also suggest truncating the strength envelope for low normal stresses as follows: τ/σ should always be smaller than 70° and, in this case, the envelope should go through the origin ($\sigma_n = \tau = 0$ MPa), in other words the cohesion is zero.

The JRC and JCS parameters are dependent on fracture length. The measured JRC_0 and $JCS₀$ values relate to fracture specimens of different lengths. Therefore, the measured values are normalised and extrapolated to values that relate to a standard fracture length of 100 mm, and hereafter referred to as JRC_{100} and JCS_{100} values.

The parameters of the Barton and Bandis's criterion are summarised in Table 3-2 for each borehole and for all the fractures. The parameters of the Barton and Bandis's for each fracture set are summarised in Table 3-3. The fracture sets are given according to the DFN model of the site /Hermansson et al. 2005/, see Table 3-1.

It can be observed that, independently on the fracture orientation and borehole, the fracture parameters do not noticeably change. Some of the tested samples were mapped as "sealed fractures" in BOREMAP maybe because of some mismatch in the reported sample depth.

Borehole	Number of samples	Φ_{h}^{BB} [°]	Φ_r^{BB} [°]	JRC_{100} [-]	JCS_{100} [MPa]
KAV01	26	30.8(0.8)	26.3(2.2)	6.2(1.6)	53.0 (13.2)
KSH01A	41	31.2(2.6)	26.2(3.1)	6.1(1.2)	76.2 (25.7)
KSH02A	48	31.5(1.6)	26.2(3.4)	5.8(1.4)	70.3 (25.7)
KLX02	24	31.4(1.2)	25.4(2.9)	6.7(1.5)	63.3 (21.4)
KLX04	18	31.4(1.0)	25.2(2.3)	5.8(1.8)	60.0 (19.4)
All fractures	157	31.3(1.7)	26.0(2.9)	6.1(1.5)	66.7 (24)

Table 3‑2. Summary of the results of tilt tests performed on rock fractures sampled from Borehole KAV01, KSH01A, KSH02, KLX02, and KLX04.

The average values are indicated. The standard deviation is set between brackets.

The average values are indicated. The standard deviation is set between brackets.

For a certain level of stresses, or for a certain stress interval, the relation in Equation (2) can be linearly approximated to determine the peak friction angle and cohesion of the Mohr-Coulomb Strength Criterion in Equation (3) (Figure 3-1):

$$
\tau = c_p^{MC} + \sigma_n \tan(\Phi_p^{MC})
$$
\n(3)

where c_p^{MC} and Φ_p^{MC} are peak cohesion and peak friction angle. Similarly, the residual cohesion and peak friction angle, c_r^{MC} respectively Φ_r^{MC} , can be fitted by the Mohr-Coulomb residual envelope. The determined Mohr-Coulomb model parameters are reported in Table 3-9 and Table 3-10.

Figure 3‑1. Fitting of the Barton Bandis' Strength Criterion with the Mohr-Coulomb's Strength Criterion.

3.2 Direct shear tests

Shear test results on natural fractures are obtained by means of three different methods. In the first method used (here referred to as Shear I), the rock specimens were clamped with cement casting with high stiffness. However, it was found that the rock specimens could slip in the cement casting. Therefore, the shear test laboratory set up was modified and epoxy was used instead as casting material. Epoxy casting reduced the problem of slipping of the rock specimen, but it has a too low stiffness. Therefore the holder device was reduced in diameter in order to avoid the effects of a soft casting material. Nineteen fracture samples from borehole KLX02 and KLX04 were tested with this modified approach at the SP laboratory (here referred to as Shear II). Furthermore, seven samples from borehole KSH01A were tested using epoxy casting at NGI (this data set is referred to as Shear III). The shear test results are summarized and analyzed in the following sections. The displacement curves were measured during normal loading and shearing so that the normal and shear stiffness of the samples could be determined. The shear tests were conducted under three levels of normal stress: 0.5, 5 and 20 MPa.

3.2.1 SP shear test results, cement casting (Shear I)

A set of 28 fracture samples from boreholes KSH01A, KSH02A and KAV01 were tested at the SP Laboratory /Jacobsson 2005opq/ based on ISRM standard from 1974. The specimens in this data set include granite to quartz monzodiorite (Ävrö granite), fine-grained dioritoid and quartz monzonite to monzodiorite.

In this shear test, the rock specimens were clamped into the specimen holder apparatus using a cement casting, which has a high stiffness. Owing to the high stiffness of the casting material, large holders could be used (a diameter of 152 mm). The benefit of large holders is that fracture core samples of different sizes and shapes can more easily be clamped into the device (i.e. with less cutting and trimming of the core specimens). Consequently, most of the available fracture area of a core sample could be shear tested by this approach (the fracture areas tested range from 20 to 52 cm²). The drawback of this casting material is that that the rock specimens could slip during the shear test. Therefore other casting materials were used for comparison.

3.2.2 SP shear test results, epoxy casting (Shear II)

Another set of 19 fracture samples from boreholes KLX02 and KLX04 were also tested at the SP laboratory /Jacobsson 2004mn/ based on ISRM standard from 1974. The specimens in this data set exclusively comprise granite to quartz monzodiorite (Ävrö granite).

In these tests, the specimens were cast with a two-component epoxy that was mixed with quartz sand to increase its stiffness. The stiffness of the epoxy mix is lower than that of cement and therefore the holder size had to be reduced to a diameter of 80 mm. Consequently, the core specimens had to be cut to fit this smaller holder, such that the tested fracture areas range from 20 to 28 cm².

3.2.3 NGI shear test results, epoxy casting (Shear III)

A set of 7 fracture samples from borehole KSH01A was tested at the NGI Laboratory /Chryssanthakis 2004e/ based on ISRM standard for shear 7testing from 1981. The specimens in this data set include fine-grained dioritoid and quartz monzonite to monzodiorite.

In these tests, the specimens were cast with epoxy that was mixed with dolomite powder. For these tests, holders with a diameter of 80 mm had to be used and therefore the core specimens had to be cut, leading to reduced tested fracture areas.

3.3 Evaluation of the mechanical parameters

Deformability and stiffness of the fractures are parameters of concern and are presented in the following sections, where results from different testing techniques are compared.

3.4 Deformability

3.4.1 Stiffness

Before shearing, the samples were normally loaded to determine their normal stiffness. The secant normal stiffness of the fracture samples for normal stress between 0.5 and 10 MPa was evaluated for the second loading cycle. The shear stiffness was then determined as the secant stiffness between 0 MPa and half of the peak shear stress. Table 3-4 shows the summary statistics for the normal and shear stiffness obtained from the tests. Most of the fracture normal stiffness data from Shear I and II range between 50 and 250 MPa/mm, but also includes exceptionally high values (primarily Shear III data). Thus, if considered all together, the normal stiffness data appears to be skewed (e.g. somewhat lognormally distributed; see Appendix 2). In such cases the arithmetic mean is not very representative of a data set (Figure 3-2). Therefore, the arithmetic mean, median and geometric mean are reported in Table 3-5. For fracture stiffness, the geometric mean is closer to the median and, therefore, the geometric mean was considered a better representation of the data.

Shear I method: data from boreholes KAV01, KSH01A, and KSH02. Shear II method: data from boreholes KLX02 and KLX04.

Shear III method: data from borehole KSH01A.

Table 3‑5. Arithmetic mean, median and geometric mean normal and shear stiffness for all the fracture samples tested with different methods.

Shear I method: data from boreholes KAV01, KSH01A, and KSH02.

Shear II method: data from boreholes KLX02 and KLX04.

Shear III method: data from borehole KSH01A.

Figure 3‑2. Frequency distributions of the normal stiffness of the fractures tested according to method Shear I, Shear II and Shear III.

3.4.2 Dilation

The dilation angle was evaluated from the tests for Shear I and Shear II (see also Appendix 2). Table 3-6, Table 3-7 and Table 3-8 show the summary statistics for the dilation angle for shear tests with normal stress 0.5, 5.0 and 20.0 MPa.

Table 3‑6. Dilation angle at normal stress 0.5 MPa evaluated for all tests in Shear I and Shear II.

Method	Number of samples	Minimum dilation angle [°]	Mean dilation angle [°]	Maximum dilation angle [°]	Standard deviation of dilation angle [°]	
Shear I	28	6.7	15.6	29.4	5.5	
Shear II	19	3.6	15.9	28.9	7.6	

Table 3‑7. Dilation angle at normal stress 5 MPa evaluated for all tests in Shear I and Shear II.

3.5 Strength

The mechanical properties of the fractures obtained from the tilt tests were already presented in Section 3.1.

In the direct shear tests, the strength envelopes of the natural fractures were rather linear so that they suited the fitting with the Coulomb's Criterion (e.g. Figure 3-3). Furthermore, peak and residual conditions could be considered. In dry conditions, the average peak and residual friction angle of all the samples were 33.8 and 29.5° , respectively. The average cohesion of the samples was 0.54 and 0.41 MPa in peak and residual conditions, respectively. Table 3-9 and Table 3-10 summarise the experimental results in terms of minimum, mean and maximum cohesion and friction angles.

The statistical parameters obtained from the three different laboratory test techniques on natural fractures are compared in Table 3-9. Distributions of the data are shown as histograms in Appendix 2.

Figure 3‑3. Peak shear and residual strength according to the Coulomb's Criterion for fracture sample KLX02-117-1.

Table 3‑9. Comparison of the results obtained from different testing techniques and laboratories: cohesion and friction angle of the peak envelope of Coulomb's Criterion. All tests samples are considered (boreholes KAV01, KSH01A, KSH02A, KLX02, and KLX04).

Laboratory test method	Number of samples	Mean		Standard deviation		Minimum		Maximum	
		$C_{\rm p}^{\rm MC}$ [MPa]	$\Phi_{p}{}^{\text{MC}}$ [°]	$C_{\rm p}^{\rm MC}$ [MPa]	$\Phi_{p}{}^{\text{MC}}$ ľΊ	C_{D}^{MC} [MPa]	$\Phi_{p}{}^{\text{MC}}$ [°]	$C_{\rm p}^{\rm MC}$ [MPa]	Φ_{p}^{MC} [°]
Tilt test ¹⁾ (KAV01, KSH01A, KSH02A, KLX02, and KLX04)	157	0.48	33.7	0.13	1.8	0.24	31.5	0.75	35.8
Shear I method (KAV01, KSH01A, and KSH02)	28	0.51	32	0.35	4.2	0.07	23.9	1.66	40.7
Shear II method (KLX02 and KLX04)	19	0.82	36.6	0.37	3.0	0.26	31.2	1.56	40.8
Shear III method (KSH01A)	7	1.1	35.4	0.18	3.8	0.89	30.2	1.36	40.3

 $¹⁾$ The values for tilt test are obtained from the Barton-Bandis' Criterion for normal stresses between</sup> 0.5 and 20 MPa.

Table 3‑10. Comparison of the results obtained from different testing techniques and laboratories: cohesion and friction angle of the residual envelope of Coulomb's Criterion. All tests samples are considered (boreholes KAV01, KSH01A, KSH02A, KLX02, and KLX04).

 $¹⁾$ The values for tilt test are obtained from the Barton-Bandis' Criterion for normal stresses between</sup> 0.5 and 20 MPa.

3.5.1 Correlation between friction angle and cohesion

The possibility of correlation between the peak cohesion and friction angle of the natural fractures was evaluated for the different sets of laboratory results (Figure 3-4). The Simpevarp and Laxemar data obtained from the tilt test and the shear test show no correlation between the peak friction angle and the peak cohesion.

Figure 3‑4. Correlations between the peak friction angle and cohesion obtained with different testing techniques and sample sets.

3.6 Statistical inference tests on fracture data

The strength and deformability data comprise several rock mechanics parameters. These have been derived from three different laboratory test methods, which may entail systematic differences in results. The primary interest is to analyze if different fracture sets have significant differences in fracture strength and deformability. The objective is therefore to distinguish whether significant differences can be found in laboratory data, in terms of: 1) fracture sets, and 2) laboratory test methods.

The parameters studied are the Mohr-Coulomb model parameters $(c_p^{MC}, \Phi_p^{MC}, c_r^{MC})$ Φ_r^{MC} and the Barton-Bandis model parameters ($\Phi_b^{\text{BB}}, \Phi_r^{\text{BB}}, \text{JRC}_{100}, \text{JCS}_{100}$). To clarify the statistical analyses, the data is referred to as x_{ij} , where x is the parameter studied, i is laboratory testing method used (Tilt, Shear I, and Shear II methods; see Section 3.1 and 3.2), and *j* is its fracture set $(S_A, S_B, S_C, S_d, \text{or } S_f)$. The fracture sets are defined in Tables 6-1 and 6-2 /Hermansson et al. 2005/. In this study, sets S_e (in Simpevarp) and S_f (in Laxemar) have been combined into a set 'S_ef', because they have similar orientation.

The total number of samples is 211 and most of the data come from tilt tests (Table 3-11). The chance of finding significant differences in statistical tests improves with larger sample sizes. However, grouping the data that do not belong to a homogeneous population may lead to incorrect inference.

The four different laboratory test methods appear to comprise rather unbiased samples of the different fracture sets (Figure 3-5). This signifies that the data can be grouped and compared, either in terms of laboratory test methods, or by fracture sets, irrespectively of risking large sampling bias. However, the modified shear method (Shear II) involves somewhat more S_A fractures, than to S_B fractures, and the original Shear method (Shear I) is applied to relatively more S ϵ fractures than to S ϵ B fractures (Figure 3-5).

Therefore, with respect to the number of data available, the following analyses are considered possible:

- 1) Analyzing influence of laboratory tests method for all fractures, regardless of fracture sets, i.e. comparing x_i , for all *j* combined.
- 2) Analyzing differences between fracture sets for tilt test data alone.

The reason for excluding all shear test data in alternative 2) is that, if the laboratory test methods entail systematic differences in results, it may be erroneous to group different types of laboratory data as one homogeneous population. The reason for not analyzing the influence of fracture set for the two types of shear test data (as done for tilt data in alternative 2) is that much less data are available from these testing methods (Table 3-11).

1) From Boreholes KAV01, KSH01A, and KSH02 in Simpevarp, and KLX02 and KLX04 in Laxemar.

2) From Boreholes KAV01, KSH01A, and KSH02 in Simpevarp.

3) From Boreholes KLX02 and KLX04 in Laxemar.

Figure 3‑5. Frequency of the samples grouped in fracture sets. "Global" refers to the total data set, regardless of testing method.

3.6.1 Statistical tests

The following three types of parametric tests were used: i) *t*-test, for pair-wise sample mean comparison, ii) *F*-test, for pair-wise sample variance comparison, and iii) ANOVA (Analysis of variance) e.g. /Davis 2002/. These are all parametric tests, which require that the underlying distributions of data are normal. Normality can be tested by various methods, including probability plots, Chi-squared, and Kolmogorov-Smirnov tests. In this study, normality was only estimated roughly by visual inspections of the data plotted as histograms (Figure 3-6, Figure 3-7, Figure 3-8, Figure 3-9). Yet, another requirement for these tests is that the sample variances and sample sizes should not be too different from each other.

If the underlying distributions prove not to be normal, non-parametric tests must be used. Therefore, the Mann-Whitney test is also used for pair-wise sample mean comparisons of sample medians (as alternative to the *t*-test) e.g. /Davis 2002/. The drawback of many nonparametric tests is that they generally require large sample sizes to be powerful. The tests use a significance level α = 0.05 (Type I error; the risk of erroneously rejecting a true null-hypothesis). The risk of the statistical tests not being powerful enough to reveal a true significant difference between data sets (Type II error) is not addressed in this study. The inference tests are finally reported in terms of *p*-values, which are defined as the smallest level of significance at which the null hypothesis would be rejected for a specific test (Table 3-12 to Table 3-18). In other words, a low *p*-value (i.e. $p < 0.05$) means that two data sets are likely to be significantly different, in terms of the specific test used.

*t***-test: pair-wise comparison of sample means**

The t-test is used to test if there is a significant difference in mean values of two data sets. The two data sets are assumed to belong to two populations that are normally distributed and have unknown and unequal variances:

$$
x_1 \sim N(\mu_1, \sigma_1^2) \text{ and } x_2 \sim N(\mu_2, \sigma_2^2)
$$
 (4)

The population mean values are estimated by the sample means ($\mu_1 \approx \bar{x}_1$ and $\mu_2 \approx \bar{x}_1$) and population variances are estimated by sample variances ($\sigma_1 \approx s_1$ and $\sigma_2 \approx s_2$).

Two hypotheses are set up: the null-hypothesis and the alternative hypothesis:

$$
H_0: \mu_1 - \mu_2 = \delta \text{ and } H_1: \mu_1 - \mu_2 \neq \delta; \quad \delta \tag{5}
$$

The test statistics used for a two-sided t-test at significance level α = 0.05 is:

$$
t = \frac{(\bar{x}_1 - \bar{x}_2 - \delta)}{\sqrt{(s_1^2 / n_1 + s_2^2 / n_2)}}
$$
(6)

Then, the hypothesis H_{A0} can be rejected at significance level α = 0.05, if:

$$
|t| \ge \frac{w_1 t_1 + w_2 t_2}{w_1 + w_2} \tag{7}
$$

where:

$$
w_1 = \frac{s_1^2}{n_1} \text{ and } w_2 = \frac{s_2^2}{n_1} \tag{8}
$$

$$
t_1 = t_{1-\alpha/2, n_1-1} = t_{0.975, n_1-1}
$$
 and $t_2 = t_{1-\alpha/2, n_2-1} = t_{0.975, n_2-1}$ (9)

The two data sets must have 'similar' variance, if the number of data n < 30. Also, the number of available data of the two data sets should not differ too much.

*F***-test: pair-wise comparison of sample variances**

The *F*-test is used to test if there is a significant difference in variance of two data sets. As above, two hypotheses are set up: the null-hypothesis and the alternative hypothesis:

$$
H_0: \sigma_1^2 = \sigma_2^2
$$

\n
$$
H_1: \sigma_1^2 \neq \sigma_2^2
$$
\n(10)

In this case, the test statistics used for a one-sided F-test at significance level $\alpha = 0.05$ is:

$$
F_c = s_1^2 / s_2^2; \quad s_1^2 > s_2^2 \tag{11}
$$

where F_c is F-distributed with n_1 –1 and n_2 –1 degrees of freedom. The null hypothesis H₀ is therefore rejected if:

$$
F_c > F_{1-\alpha, n_1-1, n_2-1} \tag{12}
$$

ANOVA: simultaneous analysis of variance

The *t*- and *F*-tests can only infer differences between two data sets at a time. The present data set can be divided into several different groups (fracture sets and laboratory test methods). Thus, the *t*- and *F*-tests require many pair-wise test combinations, either between a global group and a sub-group, or between two sub-groups. The benefit of the ANOVA test is that it can compare sample mean values of several different sub-groups at the time. The one-way ANOVA test uses the null hypothesis that assumes that there is no difference

between any of the sub-groups being compared. Then it uses the F-test statistics to compare whether the ratio (variance between groups/variance within sample groups) exceeds its critical value for rejecting its null hypothesis. The drawback is that it can only reject or accept the null hypothesis that all sub-groups share the same mean; it cannot infer if a particular group is different.

Mann-Whitney: pair-wise comparison of sample medians

The Mann-Whitney test is a non-parametric equivalence of the t-test, which tests the hypothesis that two samples have equal medians. Thus, the method is useful when the underlying distribution of a sample cannot be assumed to be normal. The two data sets are first combined. Each data value is then ranked by its value relative to others in the combined data set. Next, the ranks of combined data set are transferred back to the two original data sets. Thus, if both data sets reflect the same underlying population, the ranks would be more or less uniformly distributed between the two data sets. The test statistics used is the sum of ranks in the smaller sample set, Wx. For data sets larger than 7, Wx tends to follow a normal distribution. In such cases, it can be converted to the standard score, z (i.e. by subtracting its expected mean and dividing by its expected standard deviation) and evaluated by a two-tailed N-test (i.e. cumulative value of a standardized normal distribution, $N(0,1)$).

3.6.2 Comparison of the different laboratory techniques on natural fractures

Statistical inference tests were used to evaluate the influence that different laboratory test techniques may have on results. The four data sets compared are: tilt data (referred to as Tilt), and the three types of shear data (referred to as Shear I, Shear II and Shear III). The fracture properties studied are primarily the Mohr-Coulomb model parameters: peak cohesion c_p , peak friction angle Φ_p , residual cohesion c_r , and residual friction angle Φ_r . In this inference test, the data are only grouped in terms of laboratory testing technique, i.e. no distinction was made regarding fracture sets ($\bar{x}_i = \sum_i x_{ij}/n_i$).

The distributions of x_i were plotted as histograms to examine normality, which is required for parametric tests (Figure 3-6; Figure 3-7). In particular, Shear I appears to be more skewed than is expected for normally distributed data (Figure 3-7). However, if all three types of shear data are combined to form a larger data set, the distributions appear much more normal (Figure 3-6). Therefore the tilt test data was first compared to the grouped Shear data. The results of the *t*-tests (H_{A0}: $\mu_{\text{Tilt}} = \mu_{\text{Shear}}$), *F*-tests (H_{B0}: $\sigma_{\text{Tilt}} = \sigma_{\text{Shear}}$) and Mann-Whitney tests yield significant differences for all parameters, with the only exception of average peak friction angle (Table 3-12). Pair-wise Mann-Whitney tests are also made between tilt test data and each of the three types of shear tests (Table 3-14). The reason for this is that, in case the three types of shear tests are significantly different (as examined below), it may be erroneous to combine them into one group.

Table 3‑12. Inference tests on laboratory test methods. Tilt test data and combined shear data.

		Average, X_i						Standard Deviation, s			
		C_n _{MC}	$\mathbf{\Phi}_{\mathsf{n}}^{\mathsf{MC}}$		$C_{\rm LMC}$ $\Phi_{\rm LMC}$ $C_{\rm LMC}$ $\Phi_{\rm C}$ $C_{\rm LMC}$				Φ ^{MC}		
Tilt test data			157 0.48 33.73 0.43 28.47 0.13 1.76 0.11 3.33								
All Shear data	54		0.69 34.03 0.38 32.40 0.40 4.34 0.23 4.28								

²⁾ significance level α = 0.01

 $3)$ significance level $α = 0.001$

Figure 3‑6. Exploring normality in parameter distributions of tilt test data and shear test data. The tilt test data appears normally distributed.

Figure 3‑7. Exploring normality in parameter distributions obtained by tilt test, Shear I, Shear II, and Shear III laboratory test methods. The different types of shear data do not appear normally distributed.

Table 3‑13. Pair-wise Mann-Whitney tests on laboratory test methods; tilt test data compared to each of the three different types of shear data.

28		
19		
		0.30 0.10 0.0^{3} 0.0^{2} (0.0^{3}) (0.0^{3}) (0.0^{2}) (0.0^{3}) 0.0^{3} 0.14 0.0^{2} 0.0 ³

²⁾ significance level α = 0.01

 $3)$ significance level $α = 0.001$

Average, X_i	n_i	C_p MC	Φ_{p}^{MC}	C_r _{MC}	Φ _r _{MC}	\mathbf{k}_n	k_{s}
Shear I	28	0.51	31.95	0.33	30.86	135.2	29.28
Shear II	19	0.82	36.61	0.36	33.99	237.3	41.37
Shear III	7	1.10	35.36	0.60	34.21	607.9	21.20
Standard Deviation, Si	n_i	C_p MC	Φ_{p}^{MC}	C_r _{MC}	Φ _r _{MC}	\mathbf{k}_n	k_{s}
Shear I	28	0.35	4.24	0.27	4.67	151.7	10.55
Shear II	19	0.37	3.01	0.12	3.26	78.68	11.63
Shear III	7	0.18	3.84	0.12	3.06	393.8	8.72
Pairwise Mann-Whitney tests, rejection risk H0: mi1 = mi2		C_p^{MC}	$\Phi_{p}{}^{MC}$	C_r _{MC}	$\Phi_r{}^{\text{MC}}$	k _n	k_{s}
Shear I vs II		0.0^{2}	0.0^{3}	0.16	0.01^{1}	0.0^{3}	0.0^{3}
Shear II vs III		0.04	0.20	0.0^{3}	0.40	0.0^{3}	0.0^{3}
Shear I vs III		0.0^{3}	0.053	0.0^{2}	0.057	0.0^{3}	0.076

Table 3‑14. Pair-wise inference tests on the three types of shear test laboratory methods: Shear I, Shear II, Shear III data.

²⁾ significance level α = 0.01

 $3)$ significance level α = 0.001

Having inferred significant differences between the tilt test data and the grouped shear data, the next step was to test if there are also significant differences between the three types of shear data. The data from the three shear tests were compared pair-wise, similarly as was done above (Table 3-12). The parameter distributions of the different types of shear data appear to deviate from normality (Figure 3-7). Therefore, only the non-parametric Mann-Whitney test was used.

The tests reveal significant differences in results from the three different laboratory methods (Table 3-14). In these analyses, it should be remembered that that the three types of shear data sets contain somewhat different proportions of set S_A, S_B, and S_ef fractures (Figure 3-5). The potential differences between fracture sets are examined below.

3.6.3 Comparing fracture set properties: Mohr-Coulomb model parameters

To avoid erroneous grouping of different laboratory data types, only the tilt test data set was used to analyse the differences between fractures of different sets. The reason for selecting the tilt test data set is that it is the largest homogeneous data set (Table 3-11). Too few data are available for pursuing similar studies on individual shear data sets. In this section, the fracture set properties are compared in terms of Mohr-Coulomb model parameters. An analogous study, in terms of Barton-Bandis model parameters, is given in the following section. As above, the distributions of $x_{\text{Tit},j}$ were plotted as histograms to examine normality (Figure 3-8).

Figure 3‑8. Exploring normality of Mohr-Coulomb model parameter distributions in the tilt test data set.

The ANOVA test is first used for simultaneous comparisons between the individual fracture sets. No significant differences are found between the fracture sets by the ANOVA-test (Table 3-15). Next, all tilt test data are combined to form a reference data set, *xall*, which contains data from all fracture sets. The properties of each individual fracture set are then compared to this reference data set, by means of *t*-tests (H_{A0} : $\mu_i = \mu_{all}$) and *F*-tests (H_{B0}: $\sigma_i = \sigma_{all}$). No significant differences are found between the fracture sets. The only exception being variance in peak friction angle (Table 3-15).

Next, the properties of the individual fracture sets were pair-wise compared, by means of *t*-tests (H₀: μ_{*il*} = μ_{*i*}₂), Mann-Whitney tests (H₀: m_{*il*} = m_{*i*}) and *F*-tests (H₀: σ_{*il*} = σ_{*i*}). A significant difference in mean cohesion was found between sets S_A and S_d (Table 3-16). Furthermore, the significant difference in peak friction angle variance (found in Table 3-15) is differentiated, in terms of which sets differ significantly from another (Table 3-16).

Figure 3‑9. Exploring normality of the Barton-Bandis model parameter distributions in the tilt test data set.

1) ANOVA test

²⁾ significance level α = 0.05.

			t-test of sample means; risk of rejecting H ₀ : $\mu_{j1} = \mu_{j2}$						
$\mathbf{c_{p}}^{\text{MC}}$	SВ	s_c	S_d	S ef	Φ_{p}^{MC}	S_B	S_C	S_d	S_ef
S_A	0.30	0.80	0.03 ¹	0.49	S_A	0.55	0.90	0.23	0.71
S B		0.48	0.43	0.84	S_B		0.71	0.93	0.42
S_C			0.11	0.66	S_C			0.56	0.71
S_d				0.37	S_d				0.17
C_r^{MC}	S B	S_C	S_d	S_{ef}	$\Phi_r{}^{\text{MC}}$	S B	S_C	S_d	S_ef
S_A	0.46	0.99	0.03 ¹	0.60	S_A	0.65	0.54	0.58	0.39
S_B		0.51	0.27	0.90	S B		0.93	0.34	0.73
S_C			0.06	0.63	S_C			0.22	0.77
S_d				0.28	S_d				0.15
						Mann-Whitney test of sample medians; risk of rejecting H ₀ : $m_{1} = m_{12}$			
$\mathbf{c}_{\mathrm{p}}^{\mathrm{MC}}$	SВ	S_C	S_d	S_ef	$\Phi_{p}{}^{\text{MC}}$	S_B	S_C	S_d	S_ef
S_A	0.95	0.65	0.03 ¹	0.12	S_A	0.66	0.90	0.23	0.66
S B		0.44	0.54	0.75	S B		0.92	0.51	0.49
S_C			0.19	0.35	S_C			0.51	0.56
S_d				0.96	S_d				0.16
C_r MC	S_B	S_C	S_d	S_ef	$\Phi_r{}^{\text{MC}}$	S_B	S_C	S_d	S_ef
S_A	0.21	0.72	0.03 ¹	0.15	S_A	0.62	0.70	0.65	0.43
S B		0.58	0.35	0.53	S B		0.90	0.35	0.73
S_C			0.14	0.29	S_C			0.33	0.78
S_d				0.91	S_d				0.33
			F-test of sample variance; risk of rejecting H ₀ : $\sigma_{1} = \sigma_{2}$						
$\mathbf{c_p}^{\text{MC}}$	S_B	S_C	S_d	S_ef	$\Phi_{p}{}^{\text{MC}}$	S_B	S_C	S_d	S_ef
S_A	0.37	0.38	0.46	0.13	S_A	0.0^{2}	0.0^{2}	0.16	0.21
S B		0.49	0.31	0.22	S B		0.41	0.02	0.02 ¹
S_C			0.33	0.22	S_C			0.01^{2}	0.03 ¹
S_d				0.07	S_d				0.49
C_r^{MC}	S B	S_C	S_d	S_ef	$\Phi_r{}^{\text{MC}}$	S_B	S_C	S_d	S_ef
S_A	0.24	0.32	0.44	0.08	S_A	0.17	0.41	0.28	0.33
S B		0.42	0.24	0.25	S_B		0.13	0.28	0.31
S_C			0.34	0.20	S_C			0.22	0.27
S_d				0.06	S_d				0.48

Table 3‑16. Inference tests on fracture set properties of the Mohr-Coulomb model: pair-wise tests between individual fracture sets, *xj1* **and** *xj2***.**

²⁾ significance level α = 0.01

3.6.4 Comparing fracture set properties: Barton-Bandis model parameters

In this section, fracture set properties are compared in terms of Barton-Bandis model parameters (analogous to the approach above). By visual inspection, only the distribution of joint compressive strength, $JCS₁₀₀$, appears to deviate particularly from normality. For consistency, it is treated with the same tests as in the previous section, although results from the parametric *t*- and *F*-tests should be treated with caution.

Significant differences in mean basic friction angle are found between sets S_A and S_d. However, the difference in median basic friction angle between sets S_A and S_d is not significant, which may imply that the non-parametric test is inappropriate in this case. Significant differences in variance of basic friction angle are also found between sets.

Table 3‑17. Inference tests on fracture set properties of the Barton-Bandis model in Equation (2): pair-wise tests between individual fracture set, *xj***, and total data set,** *xall***.**

	Average, X_i					t-test; risk of rejecting H ₀				
	n_i	$\boldsymbol{\Phi}_{\rm b}^{\rm BB}$	Φ_r^{BB}	JRC ₁₀₀	JCS_{100}	$\Phi_{\rm b}^{\rm BB}$	$\pmb{\Phi}_\mathsf{r}^{\texttt{BB}}$	JRC ₁₀₀	JCS ₁₀₀	
S_A	26	31.4	25.9	6.51	65.8	0.51	0.85	0.19	0.87	
S_B	23	30.9	26.1	6.11	71.1	0.51	0.87	1.00	0.43	
S_C	21	31.1	26.1	6.41	69.8	0.76	0.83	0.36	0.54	
S d	66	31.3	25.8	5.85	64.1	0.77	0.60	0.21	0.47	
S_ef	21	31.4	26.6	6.15	68.3	0.64	0.44	0.93	0.75	
All	157	31.3	26.0	6.11	66.7					
			ANOVA test; risk of rejecting H_0			0.77^{1}	0.86^{1}	0.291	0.73^{1}	
		Standard Deviation, S_i				t-test; risk of rejecting H ₀				
	n_i	$\boldsymbol{\Phi}_{\rm b}^{\rm BB}$	Φ ^{BB}	JRC_{100}	JCS_{100}	Φ_{h}^{BB}	$\Phi_{\rm r}^{\rm BB}$	JRC ₁₀₀	JCS_{100}	
S_A	26	1.09	2.5	1.40	26.2	0.0^{3}	0.16	0.41	0.25	
S_B	23	2.69	3.6	1.46	24.3	0.04	0.09	0.52	0.42	
S_C	21	2.21	2.7	1.37	21.1	0.05	0.35	0.37	0.27	
S_d	66	1.42	2.9	1.42	24.9	0.03^{2}	0.50	0.38	0.34	
S_ef	21	1.48	3.1	1.75	20.8	0.21	0.31	0.12	0.24	
All	157	1.73	2.9	1.5	23.9					

¹⁾ ANOVA test

²⁾ Significance level α = 0,05

 $3)$ Significance level $α = 0.01$

⁴⁾ Significance level α = 0,001

t-test of sample means; risk of rejecting H ₀ : $\mu_{j1} = \mu_{j2}$										
Φ_{b}^{BB}	S_B	S_C	S_d	S_ef	Φ_r^{BB}	S_B	S_C	S_d	S_ef	
S A	0.36	0.54	0.71	1.00	S A	0.79	0.76	0.84	0.43	
S B		0.76	0.45	0.40	S B		0.99	0.67	0.67	
S_C			0.67	0.58	S_C			0.60	0.63	
S_d				0.78	S_d				0.31	
JRC ₁₀₀	S B	S_C	S d	S_ef	JCS ₁₀₀	S B	S_C	S_d	S_ef	
S_A	0.33	0.81	0.04^{1}	0.44	S_A	0.47	0.57	0.77	0.72	
S B		0.49	0.46	0.94	S B		0.85	0.25	0.68	
S_C			0.11	0.59	S_C			0.31	0.82	
S_d				0.48	S d				0.45	
					Mann-Whitney test of sample medians; risk of rejecting H ₀ : $m_{1} = m_{12}$					
Φ_{b}^{BB}	S_B	S_C	S_d	S_ef	$\Phi_r^{\;\;\text{BB}}$	S_B	s_c	S_d	S_ef	
S_A	0.79	0.74	0.65	0.61	S_A	0.55	0.86	0.07	0.41	
S B		0.79	0.93	0.39	S_B		0.93	0.36	0.71	
S_C			1.00	0.56	S_C			0.58	0.56	
S_d				0.41	S_d				0.27	
JRC ₁₀₀	S_B	s_c	S_d	S_ef	JCS ₁₀₀	S_B	s_c	S_d	S_ef	
S _A	0.31	0.86	0.07	0.10	S _A	0.35	0.40	0.89	0.34	
S B		0.47	0.55	0.48	S B		0.98	0.19	0.85	
S_C			0.18	0.23	S_C			0.23	0.99	
S d				0.75	S_d				0.32	
					F-test of sample variance; risk of rejecting H ₀ : $\sigma_{1} = \sigma_{2}$					
Φ_{b}^{BB}	S B	S_C	S_d	S_ef	Φ_r^{BB}	SВ	S_C	S_d	S_ef	
S_A	0.03	0.03	0.07	0.07	S _A	0.04^{1}	0.34	0.18	0.13	
S_B		0.19	0.03	$0.0**$	S B		0.11	0.11	0.29	
S_C			0.0^{2}	0.04	S_C			0.36	0.25	
S_d				0.38	S_d				0.32	
JRC ₁₀₀	S_B	S_C	S_d	S_ef	JCS_{100}	S_B	S_C	S_d	S_ef	
S_A	0.92	0.46	0.49	0.14	S_A	0.37	0.16	0.36	0.15	
S _B		0.39	0.41	0.21	S_B		0.26	0.47	0.24	
S_C			0.45	0.14	S_C			0.21	0.47	
S_d				0.10	S_d				0.18	

Table 3‑18. Inference tests on fracture set properties of the Barton-Bandis model; pair-wise tests between individual fracture sets, *xj1* **and** *xj2***.**

²⁾ significance level α = 0.01

³⁾ significance level α = 0.001

3.7 Discussion

Natural fractures were tested according to four different techniques (tilt tests and Shear I, II and III) and by two laboratories (SP and NGI). Tilt tests show that the average JRC_{100} of the fractures is on average around 6 while the basic friction angle is around 31°. The joint wall strength of the fractures $JCS₁₀₀$ seems to be rather high (around 65 MPa).

The normal stiffness of the fractures was obtained from the normal loading tests. However, the stiffness obtained from the NGI Laboratory results is much higher than that obtained by the SP Laboratory. The average NGI result is 607 MPa/mm, while the SP average results are 135 and 237 MPa/mm, for Shear I and Shear II, respectively. These differences, even tough the natural variability of the experimental results, indicates significant differences in results from different testing procedures at the two laboratories.

The direct shear also returned the strength of the natural fractures. The average peak cohesion of all the samples tested by the SP Laboratory was 0.5 and 0.8 MPa for Shear I and Shear II, respectively, and 32° and 37° respectively for peak friction angle. The same parameters determined from the NGI shear tests were slightly higher and equal to 1.1 MPa and 35°. However, these figures do not consider the fact that the fractures in each test batch (Shear I, II and III) belong to different fracture sets and thus should not be grouped disregarding the orientation.

When the tilt tests were re-analysed in terms of peak cohesion and friction angle, the values of 0.5 MPa and 34°, respectively, were obtained. The tilt test data set was also examined in terms of fracture set properties. The overall conclusion is that the differences between fracture sets are found to be small.

From the shear tests (Shear I, II and III), also the shear stiffness at half the peak shear strength could be determined. Average values of 29–41 MPa/mm and 22 MPa/mm were respectively obtained from the SP and NGI Laboratory, respectively.

4 Conclusions

This report contain a summary of all primary data on intact rock and fractures available for rock mechanics analyses at the time of the data freeze for Laxemar Site Descriptive Model version 1.2 (31st October 2004). For intact rock, only samples of granite to quartz monzodiorite (Ävrö granite) from borehole KLX02 and KLX04 were added to the data set. For the rock fractures, samples from borehole KLX02 and KLX04 were added to the dataset.

Thus, for the rock types quartz monzonite to monzodiorite and fine-grained dioritoid, the same mechanical properties as for Simpevarp SDM version 1.2 apply /SKB 2005/. For the samples of granite to quartz monzodiorite, differences were observed between the results from samples in borehole KLX02 and KLX04. In general, the mechanical properties of the samples from the second borehole are lower (e.g. 11% for the uniaxial compressive strength).

The samples containing sealed fractures and tested in uniaxial and triaxial conditions exhibit a Young's modulus higher than for the samples without any fractures.

Testing of fractures involved tilt testing and direct shear testing. For fracture testing, new samples were collected on all fracture sets. Therefore, the statistics of the mechanical properties for each fracture set could be updated.

Because the same mechanical parameters could be determined by means of two different testing methods, and for the same method, with different sample preparation techniques (i.e. Tilt, Shear I, Shear II, Shear III described in this report), the comparison of the different sets of laboratory results was carried out.

From the tilt tests of the fractures, it was inferred that all the fracture sets seem to have the same mechanical properties independently of the set orientation. The only significant difference was observed for the basic friction angle between set S_A and S_d. The average peak friction angle for the Coulomb's Criterion of the fracture sets varies between 33.6° and 34.1°, while the average cohesion ranges between 0.46 and 0.52 MPa, respectively. The average of the Coulomb's residual cohesion and friction angle varies in the ranges 28.0°–29.2° and 0.40–0.45 MPa, respectively. On the other hand, the direct shear testing results were too few to enable the comparison of different fracture sets.

From the direct shear tests of the fractures, it was observed that differences could be observed between the statistics of a certain parameter (e.g. peak cohesion) obtained by means of the three different sample preparation techniques. Furthermore, significant differences were observed between the parameters obtained from the tilt tests and the same parameters obtained from the direct shear tests.

Major differences were obtained for the normal stiffness of the fractures determined by NGI Laboratory (Shear III) compared to the results obtained by SP Laboratory (Shear I and II). The fracture stiffness for normal stresses between 0.5 and 10 MPa ranges between 311 and 1,461 MPa/mm for Shear III method, while it varies between 49 and 864 MPa/mm for Shear I and II methods.

Recent laboratory experiments on samples from Laxemar (KLX03A and KLX06A) also show that the fracture stiffness depend on the testing technique due to a new testing technique. This testing technique is gaining credit /Jacobsson and Flansbjer 2005ab/ and gives fracture stiffness results very similar to the Shear III results. This implies higher stiffness values than those obtained from Shear I and II methods reported in this report.

The large differences between results obtained from laboratory test methods rise questions about the comparability of the laboratory tests. Therefore, considering the uncertainty in laboratory tests, the differences found between fracture sets appear to be less relevant. Among these four data sets, the SP results were chosen to represent the properties of the fractures at Oskarshamn, mainly because they were the most numerous and agreed well with the tilt tests. In particular, SP results for the fractures from borehole KLX02 and KLX04 would be representative for the Laxemar subarea.

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Intact rock

A.1.1 Variation of the uniaxial compressive strength of the granite to quartz monzodiorite with depth

A.1.2 Variation of the Young's modulus of the granite to quartz monzodiorite with depth

A.1.3 Variation of the Poisson's ratio of the granite to quartz monzodiorite with depth

Natural fractures

A.2.1 Laboratory test methods – Mohr-Coulomb's parameters

Laboratory test methods – Mohr-Coulomb's parameters

A.2.2 Normal loading and shear tests – fracture stiffness

Note: Linear and logarithmic binning scales compared in normal stiffness histograms.

A.2.3 Tilt tests – Mohr-Coulomb's parameters for each fracture set

Mohr-Coulomb's peak cohesion for all Tilt data. The fractures are grouped into fracture sets.

Mohr-Coulomb's peak friction angle for all Tilt data. The fractures are grouped into fracture sets.

Mohr-Coulomb's residual cohesion for all Tilt data. The fractures are grouped into fracture sets.

A.2.4 Tilt tests – Barton-Bandis' parameters for each fracture set

Barton-Bandis' basic friction angle for all Tilt data. The fractures are grouped into fracture sets.

Barton-Bandis' residual friction angle for all Tilt data. The fractures are grouped into fracture sets.

Barton-Bandis' joint roughness coefficient JRC₁₀₀ for all Tilt data. The fractures are grouped into fracture sets.

A.2.5 Tilt tests – Variation with depth of the Mohr-Coulomb's parameters

Peak and residual values for cohesion and friction angle from tilt tests. The fractures are grouped per fracture sets.

A.2.6 Shear tests – Variation with depth of the Mohr-Coulomb's parameters

Peak and residual values for cohesion and friction angle from shear tests. The data are grouped per test method.

A.2.7 Shear tests – Variation with depth of the normal and shear stiffness

Fracture normal and shear stiffness, derived from normal loading and shear tests. The data are grouped per test method.

A.2.8 Shear I (SP Laboratory, cement casting) – Coulomb's parameters and stiffness from normal load and shear tests

Mechanical properties of fractures evaluated from laboratory tests (boreholes KAV01, KSH01A, and KSH02A).

A.2.9 Shear I (SP Laboratory, cement casting) – Dilation angle from shear tests

A.2.10 Shear II (SP Laboratory, epoxy casting) – Coulomb's parameters and stiffness from normal load and shear tests

Mechanical properties of fractures evaluated from laboratory tests (borehole KLX02 and KLX04)

A.2.11 Shear II (SP Laboratory, epoxy casting) – Dilation angle from shear tests

A.2.12 Shear III (NGI Laboratory, epoxy casting) – Coulomb's parameters and stiffness from normal load and shear tests

Mechanical properties of fractures evaluated from laboratory tests (borehole KSH01A)

A.2.13 Tilt test data (NGI Laboratory) – Coulomb's parameters

Mechanical properties of fractures evaluated from laboratory tests (borehole KAV01)

Borehole	Secup (unadjusted)	Peak cohesion c_n^{MC} [MPa]	Peak friction angle φ_{D}^{MC} [°]	Residual cohesion Residual friction c_n^{MC} [MPa]	angle φ_{p}^{MC} [°]	Fracture set
KAV01	434.624	0.481	31.03	0.449	27.93	d
KAV01	520.173	0.555	34.26	0.518	31.57	A
KAV01	476.288	0.254	30.8	0.246	29.4	С
KAV01	564.191	0.463	33.21	0.42	29.12	e
KAV01	516.820	0.31	31.56	0.268	24.27	d
KAV01	602.269	0.52	33.42	0.502	32.03	A
KAV01	577.760	0.602	32.89	0.495	24.03	d
KAV01	643.283	0.75	34.26	0.66	29.29	e
KAV01	625.830	0.277	33.83	0.258	30.84	e
KAV01	716.899	0.567	34.4	0.528	31.6	d
KAV01	710.070	0.504	33.19	0.487	31.8	d

Mechanical properties of fractures evaluated from laboratory tests (borehole KLX02)

Mechanical properties of fractures evaluated from laboratory tests (borehole KLX04)

Mechanical properties of fractures evaluated from laboratory tests (borehole KSH01)

Borehole	Secup (unadjusted)	Peak cohesion Peak friction c_p ^{MC} [MPa]	angle $\varphi_p{}^{\text{MC}}$ [°]	Residual cohesion c_p^{MC} [MPa]	Residual friction angle $\varphi_{p}{}^{\text{MC}}$ [°]	Fracture set
KSH01A	115.621	0.501	25.92	0.466	21.93	B
KSH01A	107.911	0.499	33.09	0.478	31.4	B
KSH01A	122.209	0.374	27.82	0.35	24.32	C
KSH01A	165.347	0.673	32.92	0.577	26.35	С
KSH01A	164.816	0.482	34.59	0.429	29.8	B
KSH01A	165.531	0.542	32.04	0.478	26.36	d
KSH01A	178.449	0.482	31.12	0.411	23.24	e
KSH01A	221.601	0.451	34.24	0.414	30.75	A
KSH01A	216.176	0.468	34.32	0.395	26.93	d
KSH01A	201.988	0.735	38.42	0.604	31.46	B
KSH01A	259.451	0.6	37.09	0.521	31.91	d
KSH01A	261.769	0.31	34.67	0.281	30.48	d
KSH01A	262.924	0.493	35.43	0.407	27.45	d
KSH01A	317.628	0.389	33.03	0.343	27.44	d
KSH01A	304.711	0.433	36.37	0.383	31.58	d
KSH01A	317.787	0.318	32.43	0.281	26.53	d
KSH01A	349.441	0.421	32.53	0.384	28.44	d
KSH01A	385.543	0.319	34.79	0.294	31.5	C
KSH01A	376.607	0.713	34.52	0.587	26.46	d

Mechanical properties of fractures evaluated from laboratory tests (borehole KSH02A)

A.2.14 Fracture set orientations

A detailed description of the derivation process for fracture sets is presented in /Hermansson et al. 2005/. The following tables summarize the chosen SDM Laxemar 1.2 fracture orientation model. These sets are based solely on univariate Fisher spherical probability distributions, and represent the 'best fit' to observed stereonet patterns. The distribution parameters were produced by entering amalgamated data from all outcrops in a particular subarea, applying a hard-sectored set division, and recording the results.

		Mean Pole		Distribution Details		$K-S$		
Set Name	Orientation Model	Trend	Plunge	Dispersion	Relative Intensity	Number of Fractures	Statistic	% Significance
S A	Univariate Fisher	338.1	4.5	13.06	28.28%	593	0.031	55.60%
S B	Univariate Fisher	100.4	0.2	19.62	26.90%	564	0.058	10.70%
S_C	Univariate Fisher	212.9	0.9	10.46	29.47%	618	0.076	15.70%
S d	Univariate Fisher	3.3	62.1	10.13	9.63%	202	0.021	99.70%
S_f	Univariate Fisher	243	24.4	23.52	5.72%	120	0.216	Not Significant

Table A2‑1. Laxemar fracture set orientations. From /Hermansson et al. 2005/.

Table A2‑2. Simpevarp fracture set orientations. From /Hermansson et al. 2005/.

		Mean Pole		Distribution Details		K-S		
Set Name	Orientation Model	Trend	Plunge	Dispersion	Relative Intensity	Number of Fractures	Statistic	% Significance
S A	Univariate Fisher	330.3	6.1	16.8	30.33%	1,190	0.091	Not Significant
S B	Univariate Fisher	284.6	0.6	10.78	18.30%	718	0.076	0.02%
S_C	Univariate Fisher	201.8	3.7	14.6	31.12%	1,221	0.043	5.20%
S_d	Univariate Fisher	84.6	81.8	6.98	8.28%	325	0.053	6.90%
S e	Univariate Fisher	67.1	15.5	11.73	11.98%	470	0.105	0.00%